

#### GEOLOGICAL SETTING - AGASSIZ METALLOTECT

The Agassiz MetalloTECT is located in the northern belt of the Lynn Lake greenstone belt and is characterized by a unique and persistent geophysical signature (Fedikow, 1986a) and a distinctive lithological association that consists of high MgO-Ni-Cr basaltic rocks (termed "picritic basalts"), iron formation, clastic sedimentary rocks (Fedikow and Gale, 1985) and felsic volcanic rocks (Parbery, 1988). The metalloTECT has a strike length of 70 km and extends from the Spider Lake area (Fedikow, 1988b) to the Sheila Lake-Margaret Lake area (Ferreira, 1988). The Nisku, MacLellan and Rainbow Au » Ag, Pb, Zn deposits, the Dot Lake Au deposit and the Farley Lake Au deposit occur along the metalloTECT (Fedikow *et al.*, 1990).

#### GENERAL STRATIGRAPHY AND LITHOLOGY

Picritic (high MgO-Ni-Cr) basaltic rocks are the most conspicuous lithology of the Agassiz MetalloTECT. Ranges for MgO, Ni and Cr are: 10.21-18.46%, 393-1 179 ppm, and 939-2 032 ppm, respectively. Picritic basaltic rocks occur throughout the metalloTECT, but decrease in abundance eastward from the MacLellan deposit. Picritic rocks overlie, and are intercalated with, oxide, subhida and by felsic volcanic rocks. Picritic rocks weather a distinctive blue green to dark green and consist of 0.5 to 4 m thick heterolithic, monolithic and flow breccia, pillow breccia, tuff and pillowed flows (Parbery and Fedikow, 1987). Locally, up to 30%, disseminated, subhedral to euhedral magnetite occurs in the picritic volcanic rocks. The rocks of the Agassiz MetalloTECT are bound to the south and north by aluminous basaltic fragmental rocks (Fedikow, 1988b) and minor felsic and mafic intrusions.

Heterolithic breccia contains two to five clast types that in total make up 40 to 80% of the rock. Clasts are commonly subrounded to subangular, amygdaloidal, and may be aphyritic or feldspar- and/or amphibole-phyric. The clasts range in size from 0.5 by 1 cm to 15 by 30 cm. Breccia groundmass is very fine grained and consists almost entirely of chlorite and amphibole with accessory magnetite. Outcrops of fragmental picritic rock that have been strongly foliated may contain up to 10%, 1 to 4 mm amphibole porphyroblasts in the matrix and clasts.

Monolithic breccia contains 20 to 50% light green clasts in a very fine grained, dark green, chlorite-amphibole matrix. Clasts are subangular to subrounded, aphyritic to very fine grained, and contain up to 10%, <2 mm plagioclase + quartz amygdaloids. Clasts are generally elongated parallel to foliation and may be a few to several centimetres in length.

Dark green basaltic rocks with distinctive higher contents of MgO, Ni and Cr (6 to 10% MgO and several hundred ppm Ni and Cr) than the aluminous basaltic rocks (average of three samples = 4.4% MgO, 57 ppm Ni and 91 ppm Cr; Syme, 1985) are intercalated with the picritic basaltic rocks. These mafic volcanic rocks occur along the length of the Agassiz MetalloTECT.

Aphyric and quartz-phyric felsic volcanic rocks occur in several locations along the Agassiz MetalloTECT and are commonly associated with picritic rocks. They are most common in the Barrington Lake area (Fedikow *et al.*, 1990). Quartz-phyric felsic rocks have been noted in drill core at the MacLellan deposit and are considered to overlie the picritic rocks (Fedikow, 1988b).

Exposures of banded iron formation (BIF) along the Agassiz MetalloTECT are sporadic. Geochemical data suggest that iron formation is present along most of the metalloTECT. Most BIF observed in the field is oxide facies iron formation, either as chert/quartz magnetite or chert/quartz-hematite. These units are generally 0.1 to 1.0 m thick, have limited extent, and are interlayered with basaltic volcanic and/or sedimentary rocks. Sulfidized magnetite/chert BIF at Farley Lake contains gold (Briggs and Taylor, 1987). Silicate facies iron formation has been observed in drill core from the MacLellan deposit and contains 5 to 20%, 5 to 10 mm pink garnets in a fine grained, green, chloritic matrix with minor magnetite and lesser amounts of calcite and amphibole. In drill core, thin 1 to 10 mm cherty layers are commonly intercalated with chlorite-rich layers. The silicate facies BIF does not appear to contain sulfides or gold. Sulfide facies iron formation occurs within picritic basalt and clastic sedimentary rocks at the MacLellan deposit. This facies of iron formation consists of 2 to 15 cm thick laminated, gold-bearing disseminated to solid pyrrhotite and pyrite layers that are rhythmically intercalated with botite and quartz-rich layers. Sphalerite, quartz and calcite occur as accessory minerals. Gagnon (1991) considers these iron sulfide-quartz layers to be deformed quartz veins.

Clastic sedimentary rocks are intercalated with picritic and nonpicritic volcanic rocks and have been referred to as siltstone, calcareous greywacke, and siliceous tuff (Fedikow, 1988b). Exposures are 0.01 to 2.0 m wide. The sedimentary rocks are fine grained, weather white to brown grey, and may contain up to 2%, 1 to 2 mm disseminated subhedral to euhedral magnetite crystals in a quartz-feldspar + biotite groundmass.

Laminated to bedded, reverse and normally graded siltstone has been identified at the MacLellan deposit. The siliceous and/or biotite-rich layers that host the sulfide mineralization and gold may represent either sedimentary rocks or zones of intense alteration arranged concentrically about a shear zone(s) (Fedikow, 1988b).

Other lithologies in the metalloTECT include mafic to intermediate volcanic rocks and small tonalitic, dioritic and gabbroic intrusions.

#### STRUCTURAL COMPONENT

Rocks in the northern belt of the Lynn Lake greenstone belt have undergone moderate physical deformation. Gilbert *et al.* (1980) describe the northern belt as consisting of a homoclinal, north-facing sequence of supracrustal rocks; however, Parbery (1988) notes that within the high-Mg (picritic) volcanic rocks, tops are commonly to the south as indicated by pillow tops, pillow breccia and graded bedding. Isoclinal folds probably resulted in both north and south facing top directions. Strike directions are dominantly eastward and dips are steep. Foliations trend mainly east-northeast. A persistent crenulation cleavage (at 244°/78°W), which can be measured over a distance of 5 km, occurs within the picritic basalts at the eastern end of the metalloTECT. Picritic rocks that outcrop in the MacLellan deposit area are characterized by the development of mylonitic textures, shear bands, and pseudo-tachylite.

#### EAGLE LAKE AREA - WEST

##### Picritic Basalts

Picritic rocks are exposed in a 500 m wide by 12 km long continuous zone from Eagle Lake to Gordon Lake (Parbery and Fedikow, 1987). Picritic basaltic rocks in the Eagle Lake - West area are fragmental.

##### Iron Formation

###### Oxide Facies Iron Formation

West of Eagle Lake BIF consists of interlayered siliceous, quartz- and chlorite-rich, and magnetite-rich laminae. The siliceous and quartz- and chlorite-rich layers are 0.1 to 0.5 cm thick and the magnetite layers are 0.1 to 1.0 cm thick. Rare hematite layers, up to 1 mm thick, occur in BIF in the Muskeg Lake area, 7 km west of Eagle Lake.

###### Silicate Facies Iron Formation

Silicate facies iron formation occurs in a small exposure west of Eagle Lake. The rock contains 5 to 20%, 5 to 10 mm pink garnets in a fine grained, green, chloritic matrix that contains minor magnetite and lesser amounts of calcite and amphibole.

##### Other

Aluminous heterolithic volcanic breccia and monolithic flow breccia occur to the north of the picritic rocks in the eastern part of the map area. Dacitic tuff and mafic tuff and breccia occur in the western and northern parts of the map area, respectively.

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### Eagle Lake Area - West Map # 1

