

MARGINAL NOTES

The Granville Lake region (NTS 64C) lies in the Churchill structural province and encompasses four Early Proterozoic geological domains (see map below). From north to south these are:

- (1) the southern margin of the Southern Indian domain,
- (2) the Lynn Lake domain,
- (3) the western part of the Leaf Rapids domain,
- (4) the northern margin of the Kisseynew domain.

Each domain has a unique stratigraphic succession of supracrustal rocks as indicated on the legend.

The Southern Indian domain is a gneiss belt comprising abundant intrusive rocks, paragneiss, amphibolite and migmatite. The metasedimentary rocks were derived from granitic, gneissic, siltstone turbidites (W) and are overlain by a thin unit of amphibolite or mafic tuff (A) and a thick metasedimentary formation (C). The conglomerate is interpreted to occupy a syncline along the north margin of the Lynn Lake domain. But a scarcity of facies relationships within the Southern Indian domain uncertain (Gilbert et al., 1980). The conglomerate contains tonalite cobbles that may have been derived from 1076 Ma plutons in the Lynn Lake domain, maximum age of the conglomerate is 1010 ± 150-10 Ma determined from zircons in the underlying volcanic rocks of the Lynn Lake domain (Baldwin et al., 1985). The megacrystic granites along the margin of the Chipewyan domain yield a minimum age of 1864 ± 15 Ma for the Southern Indian gneisses (Zwanzig et al., 1985). The southern part of the domain is dominated by tonalite (T), and granodiorite (G) which are older than the 1864 Ma granites.

The Lynn Lake domain comprises metamorphosed volcanic, sedimentary and plutonic rocks. It extends 130 km from west to east and lies on strike with the La Ronge domain in Saskatchewan (Sibbald et al., 1976; Lewry and Sibbald, 1977). The volcanic rocks make up the Wasewwan Group (Bateman, 1945; Campbell, 1969); rhyolites (W) at location 3 yielded a zircon age of 1910 ± 150-10 Ma (Baldwin et al., 1985) and a Rb-Sr metamorphic age of 1795 ± 35 Ma with 1-σ (Clark, 1980). The volcanic rocks have been intruded by subvolcanic plugs and were subsequently folded, faulted and later intruded by the Lynn Lake gabbro (B). The gabbro contained Cu-iron pipes which were exploited at mineral localities 9 and 10 (Pinnett, 1980). Large intradikes and felsic plutons of the Pool Lake intrusive suite (P, A, D), which are the pre-Sickle intrusions of Milligan (1960), yield an U-Pb zircon age of 1776 ± 40 Ma at localities 1 and 2 (Baldwin et al., in prep.) and a Rb-Sr age of 1855 ± 85 Ma with 1-σ (Clark, 1984). Unconformably overlying these rocks are metasediments and conglomerates of the Sickle Group (Norman, 1933; Campbell, 1969). Regional upper greenschist to amphibolite facies metamorphism and further deformation and plutonism took place after deposition of the Sickle Group.

The Lynn Lake domain is divided into several tectonic belts separated by granitoid intrusions: the Northern Belt, the Southern Belt and smaller belts in the granitic terrane to the south (Gilbert et al., 1980). Each belt is characterized by a unique stratigraphic succession, chemistry and structure, correlations of units between the belts are uncertain. The Southern Belt extends from Fox Mine to the area east of Barrington Lake and consists of lens-shaped volcanic and sedimentary units which have been interpreted as overlapping volcanic edifices with flanking aprons of volcanoclastic rocks. The volcanic units contain tholeiitic basalt (M) at Cookman Lake and porphyritic, "ornithine" calc-alkaline basalt (Wp) at McVeigh Lake and Fox Mine. The basalts are overlain locally by felsic volcanic bodies. Fox Mine is situated on the margin of one of the belts. At Hughes Lake the tholeiitic basalt is overlain by a predominantly andesitic calc-alkaline suite (W) (Syme, 1985). The units in the Southern Belt form a large anticline and probably contain the oldest rocks in the Lynn Lake domain (Gilbert et al., 1980). In contrast, the structure in the Northern Belt is north-facing homocline. The stratigraphic sequence of the Northern Belt has 1910 ± 10 Ma old tholeiite (Baldwin et al., in prep.) at the base, overlain in ascending order by andesite and basalt, sedimentary rocks (w, wW) and an upper basaltic section (wW). The mafic rocks are tholeiites which include some high-alumina varieties and subordinate high-magnesian andesites. The upper mafic section contains the MacLellan and Farley Lake gold deposits (mineral localities 15 and 16). The Northern Belt succession is tentatively correlated with the upper part of the Southern Belt section and thus may represent the upper part of the Wasewwan Group.

A thick sequence of metasediments (S) (Sickle Group) with basal conglomerate (C) unconformably overlies the Pool Lake intrusions and the earlier deformed rocks of the Wasewwan Group. The angular unconformity is clearly evident on the northwest shore of Sickle Lake, where southeast-trending Wasewwan Group volcanic units are abruptly truncated by the northeast-trending Sickle conglomerate. At Hughes Lake the base of the Sickle Group cuts across the volcanic stratigraphy; a major fold closure in the Sickle Group southward of the lake has no expression in the underlying Wasewwan Group rocks which form a south-facing homocline. The basal part of the Sickle conglomerate typically contains clasts derived from the directly underlying bedrock. The remainder of the conglomerate contains predominantly tholeiitic, andesite and rare granitic clasts. These tholeiites are not common in the underlying bedrock and are of uncertain provenance. The Sickle Group has characteristics of alluvial deposits; a thick basal conglomerate is overlain by a thick, fine-grained sandstone succession, trough crossbedding, ripple lamination, mudcracks and mudclasts are locally abundant in the otherwise massively bedded sandstone.

Throughout the Lynn Lake domain, various phases of deformation have affected units of different ages. The following phases (from oldest to youngest) have been recognized:

- (1) block faulting during development of Wasewwan volcanic edifices and tectono-sedimentary basins at 1.9 Ga;
- (2) large-scale east-northeast-trending upright folding (e.g. Southern Belt anticline which is truncated by the 1976 Ma Pool Lake intrusive suite);
- (3) northeast-trending folding and faulting, during and after emplacement of Pool Lake intrusions, followed by deposition of the Sickle Group and intrusion of high-level sills of Black Trout diorite (wD);
- (4) major north-trending folding and intrusion of younger granitic rocks;
- (5) easterly trending folding producing a type I interference pattern (Ramsey, 1967) well developed between Conglomerate Lake and Hunter Lake;
- (6) development of northeast- and north-trending folds and faults.

The late deformation (phases 4, 5 and 6) affected the entire Granville Lake region and was accompanied by regional metamorphism. The late deformation produced northeast-trending, faulted and steeply plunging, reactivated folds in the Wasewwan Group. The most prominent fault zone is the east-trending Johnson shear, which separates the Southern Belt from the granitoid terrane to the south. The Tod Lake-Dunphy Lakes fault system extends from the Lynn Lake domain southwest into the Kisseynew domain.

The Leaf Rapids domain lies east of Beausage Lake. This domain is a predominantly intrusive terrane that contains the 1876 Ma Rusty Lake metavolcanic belt (NTS 64B). The western part of the domain comprises the Outlaw Bay tonalite and the Eden intrusive suite. Supracrustal rocks of uncertain age occur as reefs and narrow belts within the intrusive complexes; the supracrustal rocks may include Wasewwan Group amphibolite and paragneiss in the northwest, and possibly younger rocks farther south. Metarhyolite (W) of the Wasewwan Group (Milligan, 1960) and possibly older rocks extend from Black Trout Lake to Beausage Lake and eastward to the Rusty Lake belt. These supracrustal rocks include a unique suite of felsic volcanic sandstones, Jasper pebble conglomerate and iron formation. At the contact with the Lynn Lake domain these units are unconformably overlain by the Sickle Group which contains clasts of the underlying iron formation and abundant tholeiitic clasts (Zwanzig, 1981).

The Outlaw Bay tonalite is a strongly foliated complex, possibly correlative with the 1876 Ma Pool Lake intrusive suite in the Lynn Lake domain. The Eden intrusive suite consists of younger megacrystic granites (G), biotite granite (GB) and aegirine-augite monzonite (Z) which are similar to, and possibly coeval with the 1864 Ma granites in the Chipewyan domain. The Eden intrusive suite may extend northeast into the Blackfoot batholith. Black Trout diorite dikes (wD) occur at the western margin of the Leaf Rapids domain. This diorite-quartz diorite suite has a transitional to alkaline chemistry. The younger granites of the diorite dikes.

The western part of the Leaf Rapids domain is interpreted as an upthrust structural block. The metachronous grade is middle to upper amphibolite facies. Structures include phases 4, 5 and 6 folds and faults, which are also developed in the Lynn Lake domain. The northwest-trending structure in the supracrustal rocks northeast of Beausage Lake was probably developed during phase 3.

The Kisseynew domain is a metasedimentary gneiss belt that flanks the Lynn Lake and Leaf Rapids domains on the south. It is composed of paragneiss, amphibolite and migmatite derived from Lower Proterozoic sedimentary and subordinate volcanic strata (Clark et al., 1974). This belt has undergone upper amphibolite facies metamorphism and polyphase deformation. The supracrustal succession consists of three divisions:

- (1) the lowest division (Burnwood River Metamorphic Suite) (W) is composed of quartzfeldspathic garnet-biotite gneiss, cordierite, sillimanite and andesitic migmatite derived from granitic greywacke-mudstone turbidites (Gilbert et al., 1980);
- (2) the middle division comprises thin conformably overlying amphibolite and paragneiss units (A, C, W) derived from mafic volcanic rocks and clastic to chemical sediments;
- (3) the upper division (Sickle Metamorphic Suite) conformably overlies the middle division and is composed of metagranite (G) and quartzfeldspathic gneiss (S) = hornblende or sillimanite derived from lithic sandstone, calcareous sandstone and arkose.

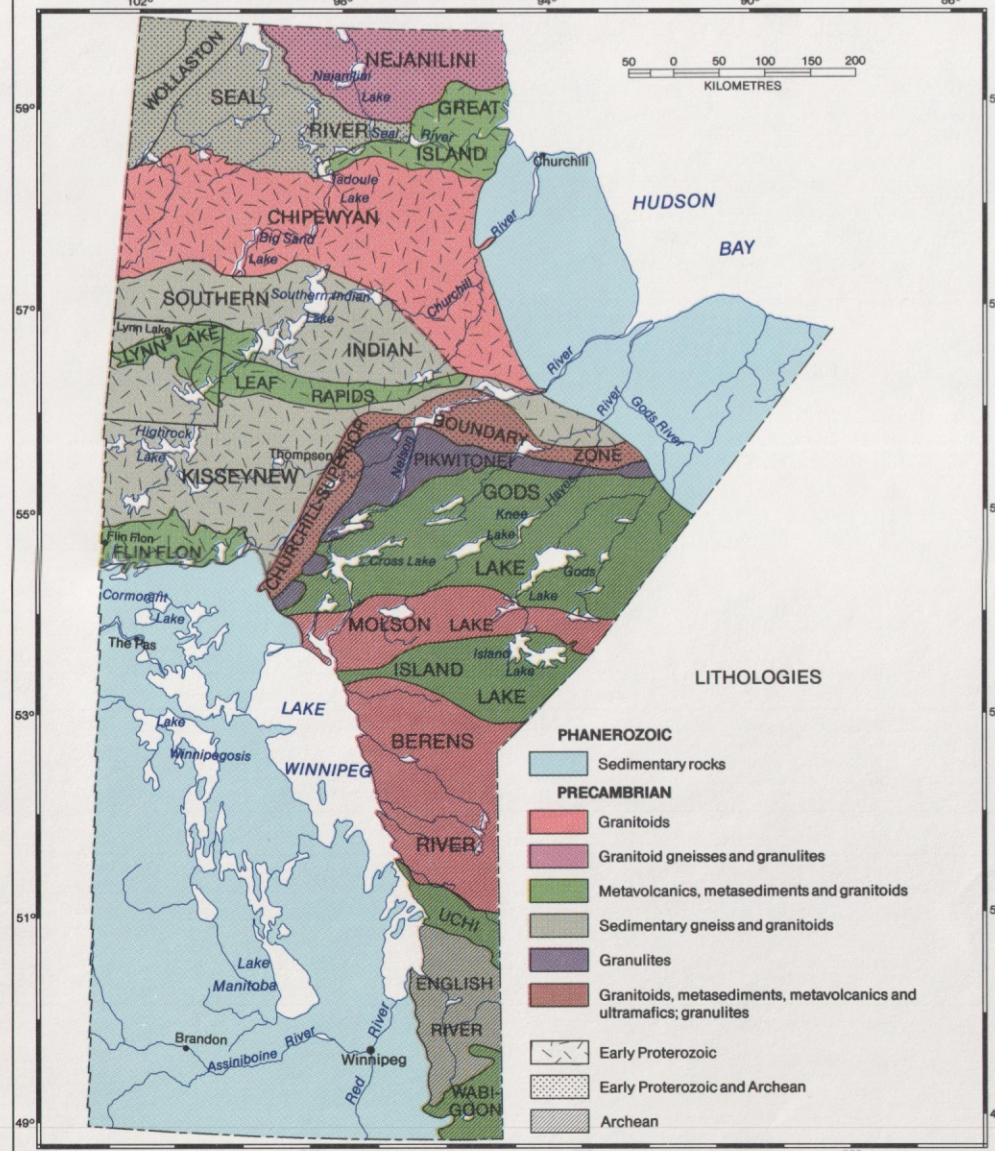
The Sickle Metamorphic Suite is stratigraphically equivalent to the Sickle Group in the Lynn Lake domain. The Burnwood River greywackes were probably deposited contemporaneously with the volcanic and sedimentary rocks of the Amiak Group in the Flin Flon domain (Bailes, 1971; McElhiney, 1974). The stratigraphic relations between the Burnwood River greywackes, the middle division and the Wasewwan Group volcanic rocks are uncertain. Rocks of the middle division occur on the north margin and in the Lynn Lake domain, northwest of the Tod Lake-Dunphy Lakes fault system. They were mapped as part of the Wasewwan Group but their contacts are generally faulted (Gilbert et al., 1980). The metasedimentary and metasedimentary rocks of the middle division are locally intercalated with the Burnwood River metagreywackes and may have the same age as the Burnwood River Suite. Feldspathic metagreywacke (W), sedimentary breccia and local metagranite (G) of the middle division were derived from high-energy turbidites with possible olistostrol facies on Granville Lake (Zwanzig, 1984). The coarse clastic rocks are associated with carbonates, metacherts, metabasalts (including rare ultramafic flows). The basalts are high-iron and high-magnesian tholeiites, geochemically transitional between modern arc-related and rift-related rocks. The middle division is stratigraphically overlain by the Sickle Metamorphic Suite but locally (structurally?) underlain by Sickle rocks.

The structure of the Kisseynew domain is dominated by large recumbent folds and upright crossfolds which trend north, east and northeast. The upright folds are coeval with late structures (phases 4, 5 and 6) in the Lynn Lake domain, and the recumbent folds probably developed between phases 3 and 4. Along the Lynn Lake-Kisseynew margin thrusting and remobilization of early plutonic suites resulted in recumbent folds and domal granitic complexes (e.g., Glasslope Lake complex).

In the Granville Lake region deposition and metamorphism of all the rocks are Early Proterozoic (Clark, 1980). The domains are similar to Archaean granite-greenstone belts and metasedimentary gneiss belts, and, like the Archaean terranes, may have evolved in a tectonic environment that has no exact modern analogue (Ayres and Thurston, 1985). However, Wasewwan Group geochemistry is similar to that of Cenozoic intra-oceanic arcs (Syme, 1985). The Pool Lake intrusive suite is calc-alkaline to calcic, and also a possible product of arc magmatism. The geochemistry and structure of the volcanic rocks at the northern margin of the Kisseynew domain (middle division) suggest fitting in an arc environment but other interpretations are possible (Zwanzig, 1984; Green et al., 1985). Synkinematic isograds (G) are slightly porphyritic and apparently had an upper crustal anastatic source, possibly related to collision tectonics.

\* Unless otherwise noted, domain descriptions are only applicable to the Granville Lake map sheet (NTS 64C).

PRINCIPAL GEOLOGICAL DOMAINS



Every possible effort has been made to ensure that the information presented on this map is accurate. However, the Province of Manitoba and Manitoba Energy and Mines do not assume liability for any errors that may occur. References are included for users wishing to verify critical information.



Scale 1:250 000  
Kilometers 0 5 10 15 20 25  
Printed 1986

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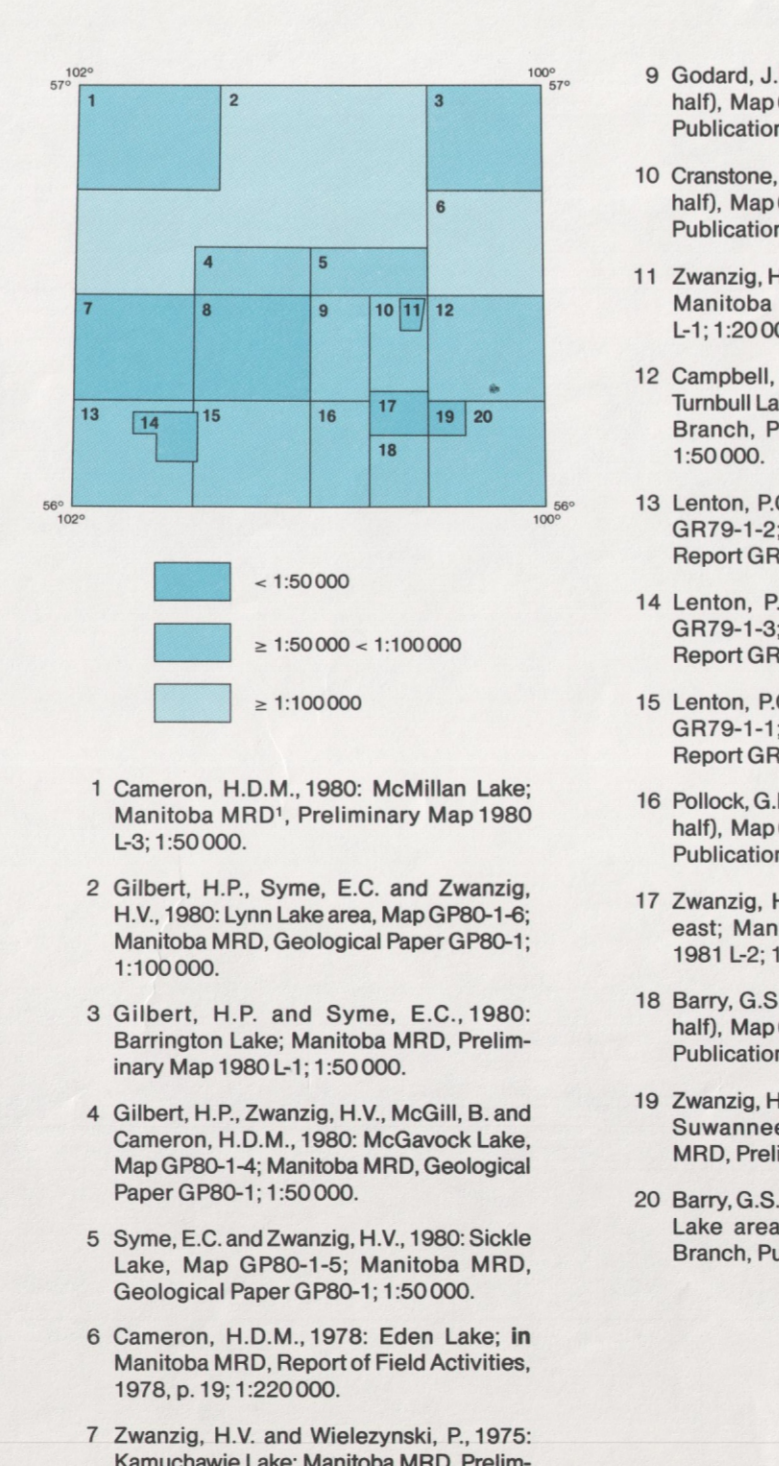
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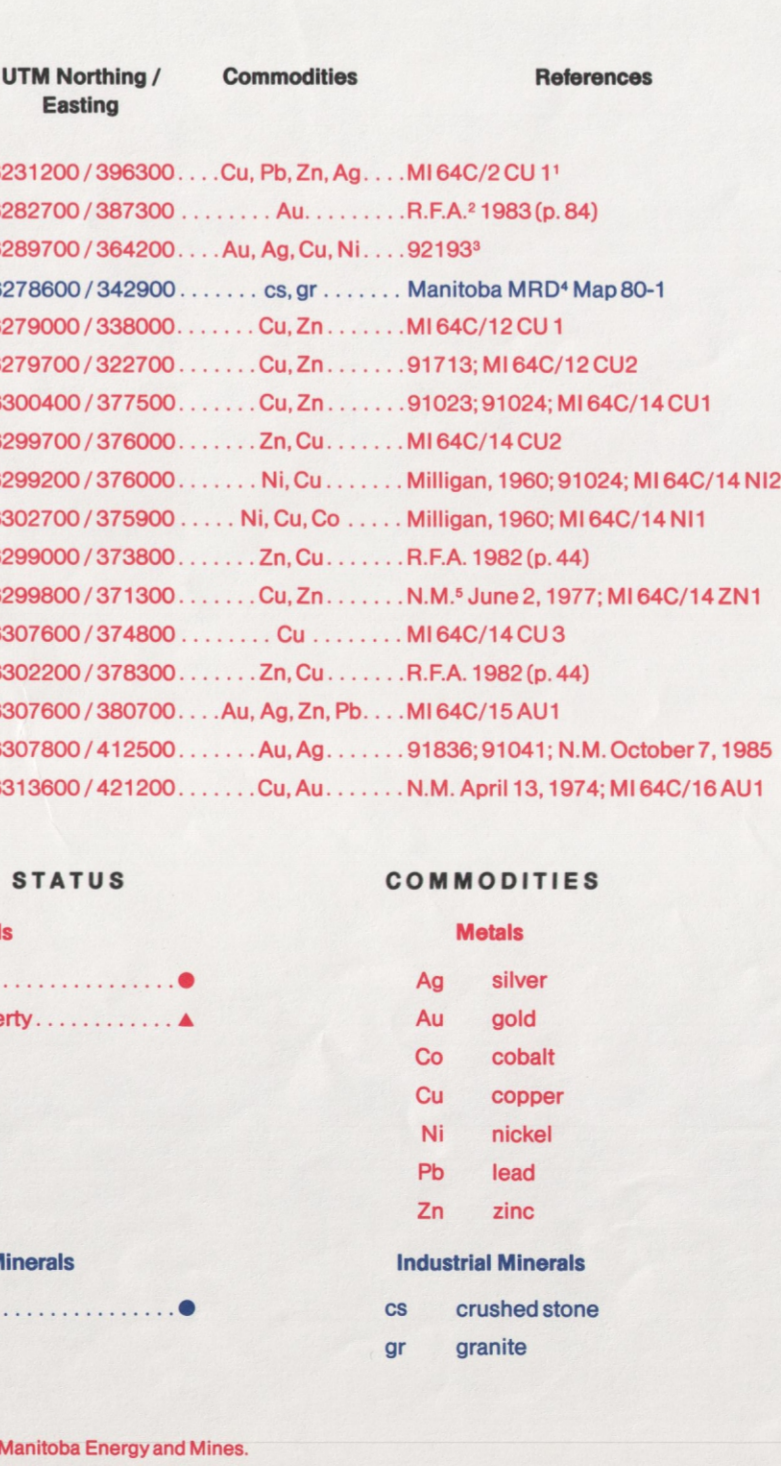
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SOURCES OF INFORMATION



LEGEND

**PRECAMBRIAN**

**Younger Plutonic Rocks**

- X Pegmatite
- Z Monzonite, monzonogranite; Z- aegirine-augite monzonite - Eden intrusive suite
- Z Syenite, syenogranite
- G Granite; Gp- megacrystic granite; Gm- magnetite granite; Gt- leucocratic granodiorite; G- granitoid (within the Glasslope Lake complex); Gb- biotite granite; Gc- megacrystic granite
- G Grandiorite; Gt- aegirine-porphyrphytic granodiorite; Gh- hornblende granodiorite; Gp- porphyroblastic granodiorite-tonalite; Gg- variolite-porphyrphytic granodiorite; Gd- granodiorite
- D Diorite, quartz diorite - Black Trout diorite
- B Gabbro

**Plutonic Rocks of Uncertain Age**

- G Granite
- G Grandiorite, granite; Gt- gneissic granodiorite; Gp- porphyroblastic granodiorite-tonalite
- T Tonalite, granodiorite; Tc- clinopyroxene tonalite; Tn- gneissic tonalite-granodiorite; Ot- Outlaw Bay tonalite; Tm- magnetite tonalite (within the Glasslope Lake complex)
- Z Monzodiorite
- D Diorite, quartz diorite; Dp- porphyroblastic quartz diorite
- B Gabbro, diorite, quartz diorite

**KISSEYNEW DOMAIN**

- SaSs Quartz-rich sillimanite gneiss derived from arkose; migmatite (patterned)
- SsB Quartzfeldspathic gneiss derived from sandstone; SsB- feldspar-rich biotite gneiss; migmatite (patterned)
- SsH Hornblende-bearing quartzfeldspathic gneiss ± diopside; migmatite (patterned)
- SsC Polymictic conglomerate; SsCh- hornblende-bearing polymictic metagranite

**LYNN LAKE DOMAIN**

- Ss Sandstone, quartzfeldspathic gneiss
- SsH Calcareous sandstone, hornblende-bearing quartzfeldspathic gneiss
- SsC Polymictic conglomerate; SsCh- hornblende-bearing polymictic metagranite

**Southern Indian Domain**

- C Polymictic conglomerate, iron formation (dotted line)
- A Amphibolite; Ar- amphibolite derived from basalt, mafic tuff and volcanic breccia
- W Greywacke, biotite gneiss, migmatite

**Leaf Rapids Domain**

- U Ultramafic and ultrabasic rocks
- A A-layered amphibolite, minor ultramafic rock, marble and felsic gneiss; Ao- amphibolite derived from pillow and massive basalt
- C Conglomerate; Cf- feldspathic sandstone-cobble conglomerate
- W Feldspathic greywacke, conglomerate, paragneiss, amphibolite; migmatite (patterned); Wt- hornblende-bearing greywacke gneiss, amphibolite, felsic gneiss
- B Gabbro, peridotite, norite, quartz diorite
- B Gabbro, diorite

**Wasewan Group (low to high metamorphic grade)**

- wV Volcanic sandstone, siltstone and mafic mudstone, iron formation (dotted line)
- wC Volcanic conglomerate, sandstone and siltstone
- wA Wasewan Group, undivided (mainly amphibolite)
- wV Rhyolite and dacite (including pyroclastic rocks) 3:1910 ± 15/-10 Ma (1790 ± 35 Ma)\*
- wV Andesite; wVr- porphyritic and aphyric basalt and andesite; wVr- intermediate
- wVr- porphyritic and aphyric basalt (undivided volcanoclastic rocks and flows), iron formation (dotted line); wVp- porphyritic basalt, massive flows and breccia; wV- mafic tuff
- Basalt; wVp- pillow basalt; wVp- basalt (mainly porphyritic); wV- aphyric and porphyritic basalt; wV- mafic tuff

**Burnwood River and/or Sickle Metamorphic Suite**

- U Ultramafic and ultrabasic rocks
- A A-layered amphibolite, minor ultramafic rock, marble and felsic gneiss; Ao- amphibolite derived from pillow and massive basalt
- C Conglomerate; Cf- feldspathic sandstone-cobble conglomerate
- W Feldspathic greywacke, conglomerate, paragneiss, amphibolite; migmatite (patterned); Wt- hornblende-bearing greywacke gneiss, amphibolite, felsic gneiss
- B Gabbro, peridotite, norite, quartz diorite
- B Gabbro, diorite

**Wasewan Group (low to high metamorphic grade)**

- wV Volcanic sandstone, siltstone and mafic mudstone, iron formation (dotted line)
- wC Volcanic conglomerate, sandstone and siltstone
- wA Wasewan Group, undivided (mainly amphibolite)
- wV Rhyolite and dacite (including pyroclastic rocks) 3:1910 ± 15/-10 Ma (1790 ± 35 Ma)\*
- wV Andesite; wVr- porphyritic and aphyric basalt and andesite; wVr- intermediate
- wVr- porphyritic and aphyric basalt (undivided volcanoclastic rocks and flows), iron formation (dotted line); wVp- porphyritic basalt, massive flows and breccia; wV- mafic tuff
- Basalt; wVp- pillow basalt; wVp- basalt (mainly porphyritic); wV- aphyric and porphyritic basalt; wV- mafic tuff

**Stratigraphic Note**

The lithologies within the major stratigraphic divisions are in approximate stratigraphic order.

LEGEND

**LEGEND**

**Geological boundary (approximate)**

**Geological boundary (inferred from aeromagnetic signature and trend)**

**Stratigraphic top (inclined, overturned)**

**Bedding, layering (dip unknown, 0°-29°, 30°-59°, 60°-79°, 80°-90°)**

**Foliation, or foliation and parallel layering (dip unknown, 0°-29°, 30°-59°, 60°-79°, 80°-90°)**

**Syncline (upright, overturned)**

**Anticline (upright, overturned)**

**Fault**

**Domain boundary**

**Area of title or outcrop**

**Sample locality for U-Pb zircon age determination**

**PROPERTY STATUS**

**Metals**

- Past producer
- Important mineral property

**COMMODITIES**

**Metals**

- Au silver
- Ag gold
- Co cobalt
- Cu copper
- Ni nickel
- Pb lead
- Zn zinc

**Industrial Minerals**

- cs crushed stone
- gr granite

**Reference for U-Pb zircon ages**

1, 2, 3, 4 Baldwin et al., 1985

\* Clark, 1984

\* Clark, 1980

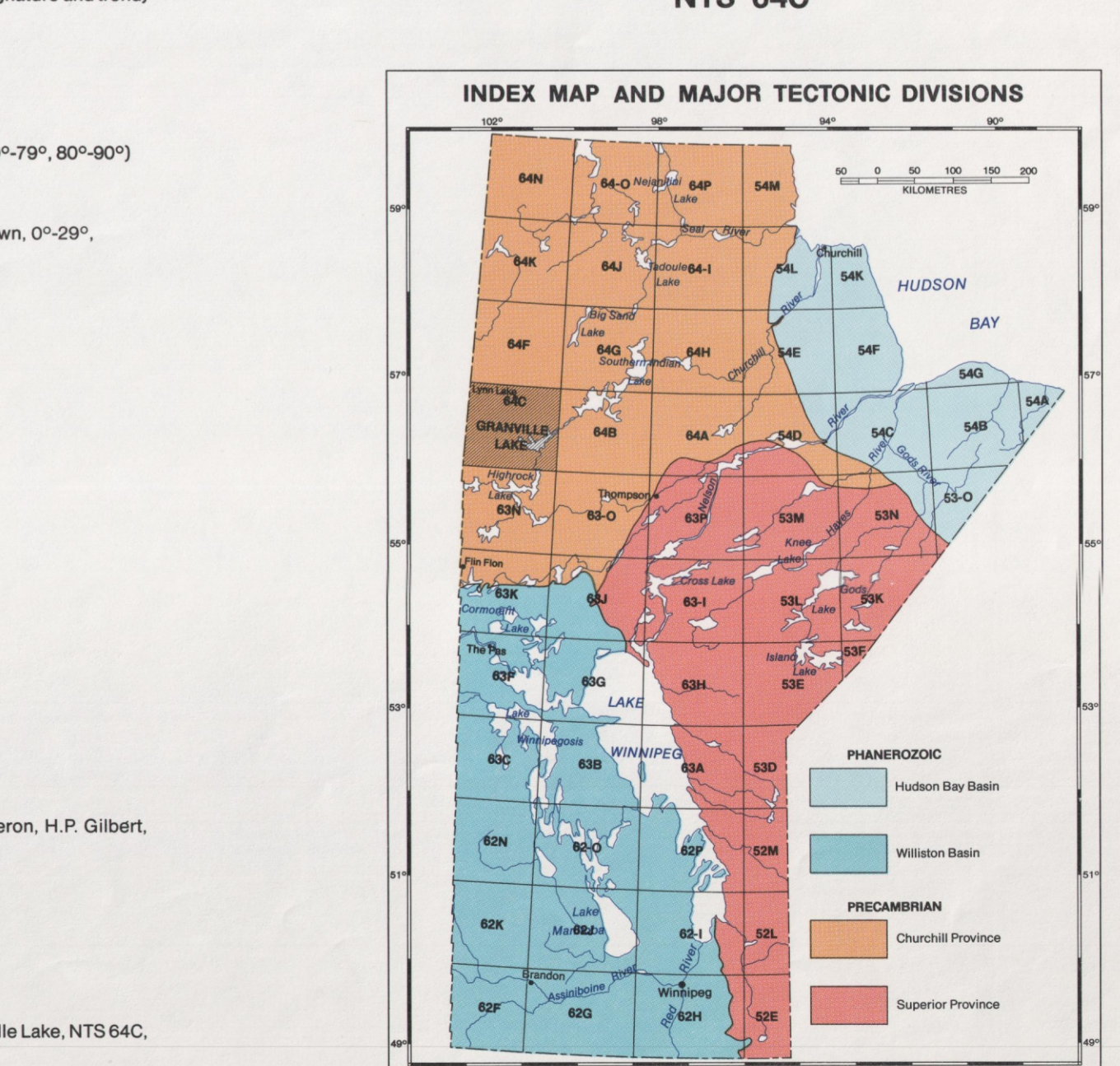
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The lithologies within the major stratigraphic divisions are in approximate stratigraphic order.

BEDROCK GEOLOGY COMPILED MAP SERIES

GRANVILLE LAKE

NTS 64C



Suggested reference to this publication: Manitoba Energy and Mines, 1986. Bedrock Geology Compilation Map Series, Granville Lake, NTS 64C, 1:250,000.

Compilation by J.S.D. Parker and D. Kowarchuk  
Cartography by D.L. McShane

A contribution by Manitoba Energy and Mines, Geological Services, to programming under the Canada-Manitoba Mineral Development Subsidy Agreement (MDA), a sub-agreement of the Economic and Regional Development Agreement (ERDA).